

## Statistical Software for Weather and Climate

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## Already familiar with R?

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Advanced (potentially useful) topics:

- Reading and Writing NetCDF file formats:  
<http://www.image.ucar.edu/Software/Netcdf/>
- A Climate Related Precipitation Example for Colorado:  
<http://www.image.ucar.edu/~nychka/FrontrangePrecip/>

## The R programming language

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R Development Core Team (2008). R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. ISBN 3-900051-07-0,  
<http://www.r-project.org>

Vance A, 2009. Data analysts captivated by R's power. *New York Times*, 6 January 2009. Available at:  
[http://www.nytimes.com/2009/01/07/technology/business-computing/07program.html?\\_r=2](http://www.nytimes.com/2009/01/07/technology/business-computing/07program.html?_r=2)

## R preliminaries

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Assuming R is installed on your computer...

In linux, unix, and Mac (terminal/xterm) the directory in which R is opened is (by default) the current working directory. In Windows (Mac GUI?), the working directory is usually in one spot, but can be changed (tricky).

Open an R workspace:

Type **R** at the command prompt (linux/unix, Mac terminal/xterm) or double click on R's icon (Windows, Mac GUI).

`getwd()` # Find out which directory is the current working directory.

## R preliminaries

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### Assigning vectors and matrices to objects:

```
# Assign a vector containing the numbers -1, 4 and 0 to an
object called 'x'
x <- c(-1, 4, 0)

# Assign a 3x2 matrix with column vectors: 2, 1, 5 and
# 3, 7, 9 to an object called 'y'.
y <- cbind(c(2, 1, 5), c(3, 7, 9))

# Write 'x' and 'y' out to the screen.
x
y
```

## R preliminaries

---

### Subsetting vectors:

```
# Look at only the 3-rd element of 'x'.
x[3]

# Look at the first two elements of 'x'.
x[1:2]

# The first and third.
x[c(1,3)]

# Everything but the second element.
x[-2]
```

## R preliminaries

---

### Saving a workspace and exiting

```
# To save a workspace without exiting R.
save.image()

# To exit R while also saving the workspace.
q("yes")

# Exit R without saving the workspace.
q("no")

# Or, interactively...
q()
```

## R preliminaries

---

### Subsetting matrices:

```
# Look at the first row of 'y'.
y[1,]

# Assign the first column of 'y' to a vector called 'y1'.
# Similarly for the 2nd column.
y1 <- y[,1]
y2 <- y[,2]

# Assign a "missing value" to the second row, first column
# element of 'y'.
y[2,1] <- NA
```

## R preliminaries

---

### Logicals and Missing Values:

```
# Do 'x' and/or 'y' have any missing values?
any( is.na( x))
any( is.na( y))

# Replace any missing values in 'y' with -999.0.
y[ is.na( y)] <- -999.0

# Which elements of 'x' are equal to 0?
x == 0
```

## R preliminaries

---

### See hierarchy of loaded packages:

```
search()
```

```
# Detach the 'SpatialExtremes' package.
detach(pos=2)
```

### See how to reference a contributed package:

```
citation("fields")
```

## R preliminaries

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### Contributed packages

```
# Install some useful packages. Need only do once.
install.packages( c("fields",      # A spatial stats package.
                  "evd",          # An EVA package.
                  "evdbayes",     # Bayesian EVA package.
                  "ismev",        # Another EVA package.
                  "maps",         # For adding maps to plots.
                  "SpatialExtremes"))

# Now load them into R. Must do for each new session.
library( fields)
library( evd)
library( evdbayes)
library( ismev)
library( SpatialExtremes)
```

## R preliminaries

---

### Simulate some random fields

```
From the help file for the fields function sim.rf
help( sim.rf)
```

```
#Simulate a Gaussian random field with an exponential
# covariance function, range parameter = 2.0 and the
# domain is  $[0,5] \times [0,5]$  evaluating the field at a  $100 \times 100$ 
# grid.
grid <- list( x= seq( 0,5,100), y= seq(0,5,100))
obj <- Exp.image.cov( grid=grid, theta=.5, setup=TRUE)
look <- sim.rf( obj)
```

```
# Now simulate another ...
look2 <- sim.rf( obj)
```

## R preliminaries

---

### Plotting the simulated fields

```
# setup the plotting device to have two plots side-by-side
set.panel(2,1)

# Image plot with a color scale.
image.plot( grid$x, grid$y, look)
title("simulated gaussian field")
image.plot( grid$x, grid$y, look2)
title("another (independent) realization ...")
```

## R preliminaries

---

### Simple plot example.

```
plot( 1:10, z <- rnorm(10), type="l", xlab="", ylab="z",
      main="Std Normal Random Sample")
points( 1:10, z, col="red", pch="s", cex=2)
lines( 1:10, rnorm(10), col="blue", lwd=2, lty=2)

# Make a standard normal qq-plot of 'z'.
qqnorm( z)

# Shut off the device.
dev.off()
```

## R preliminaries

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### Basics of plotting in R:

- First must open a device on which to plot.
  - Most plotting commands (e.g., `plot`) open a device (that you can see) if one is not already open. If a device is open, it will write over the current plot.
  - `X11()` will also open a device that you can see.
  - To create a file with the plot(s), use `postscript`, `jpeg`, `png`, or `pdf` (before calling the plotting routines. Use `dev.off()` to close the device and create the file.
- `plot` and many other plotting functions use the `par` values to define various characteristics (e.g., margins, plotting symbols, character sizes, etc.). Type `help( plot)` and `help( par)` for more information.

## Background on Extreme Value Analysis (EVA)

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### Motivation

#### Sums, averages and proportions (Normality)

- Central Limit Theorem (CLT)
- Limiting distribution of binomial distribution

#### Extremes

- Normal distribution inappropriate
- Bulk of data may be misleading
- Extremes are often rare, so often not enough data

## Background on Extreme Value Analysis (EVA)

---

### Simulations

```
# Simulate a sample of 1000 from a Unif(0,1) distribution.
U <- runif( 1000)
hist( U)

# Simulate a sample of 1000 from a N(0,1) distribution.
Z <- rnorm( 1000)
hist( Z)

# Simulate a sample of 1000 from a Gumbel distribution.
M <- rgev( 1000)
hist( M)
```

## Background on Extreme Value Analysis (EVA)

---

### Simulations

```
# Simulate maxima from samples of size 30 from
# the Unif(0,1) distribution.
Umax <- matrix( NA, 30, 1000)
for( i in 1:1000) Umax[,i] <- runif( 30)
Umax <- apply( Umax, 2, max)
hist( Umax, breaks="FD", col="blue")

# Simulate 1000 maxima from samples of size 30 from
# the Fréchet distribution.
Fmax <- matrix( NA, 30, 1000)
for( i in 1:1000) Fmax[,i] <- rgev( 30,
                                loc=2, scale=2.5, shape=0.5)
Fmax <- apply( Fmax, 2, max, finite=TRUE, na.rm=TRUE)
hist( Fmax, breaks="FD", col="blue")
```

## Background on Extreme Value Analysis (EVA)

---

### Simulations

```
# Simulate 1000 maxima from samples of size 30 from
# the normal distribution.
Zmax <- matrix( NA, 30, 1000)
dim( Zmax)
for( i in 1:1000) Zmax[,i] <- rnorm( 30)
Zmax <- apply( Zmax, 2, max)
dim( Zmax)
class( Zmax)
length( Zmax)
class( Zmax)
hist( Zmax, breaks="FD", col="blue")
```

## Background on Extreme Value Analysis (EVA)

---

### Simulations Last one

```
# Simulate 1000 maxima from samples of size 30 from
# the exponential distribution.
Emax <- matrix( NA, 30, 1000)
for( i in 1:1000) Emax[,i] <- rexp( 30)
Emax <- apply( Emax, 2, max)

# This time, compare with samples from the
# exponential distribution.
par( mfrow=c(1,2))
hist( Emax, breaks="FD", col="blue")
hist( rexp(1000), breaks="FD", col="blue")
```

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

Let  $X_1, \dots, X_n$  be a sequence of independent and identically distributed (iid) random variables with common distribution function,  $F$ . Want to know the distribution of

$$M_n = \max\{X_1, \dots, X_n\}.$$

**Example:**  $X_1, \dots, X_n$  could represent hourly precipitation, daily ozone concentrations, daily average temperature, etc. Interest for now is in maxima of these processes over particular blocks of time.

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

Could try to derive the distribution for  $M_n$  exactly for all  $n$  as follows.

$$\begin{aligned} \Pr\{M_n \leq z\} &= \Pr\{X_1 \leq z, \dots, X_n \leq z\} \\ &\stackrel{\text{indep.}}{=} \Pr\{X_1 \leq z\} \times \dots \times \Pr\{X_n \leq z\} \\ &\stackrel{\text{ident. dist.}}{=} \{F(z)\}^n. \end{aligned}$$

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

If interest is in the minimum over blocks of data (e.g., monthly minimum temperature), then note that

$$\min\{X_1, \dots, X_n\} = -\max\{-X_1, \dots, -X_n\}$$

Therefore, we can focus on the maxima.

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

Could try to derive the distribution for  $M_n$  exactly for all  $n$  as follows.

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**But!** If  $F$  is not known, this is not very helpful because small discrepancies in the estimate of  $F$  can lead to large discrepancies for  $F^n$ .

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

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$$\begin{aligned}\Pr\{M_n \leq z\} &= \Pr\{X_1 \leq z, \dots, X_n \leq z\} \\ &\stackrel{\text{indep.}}{=} \Pr\{X_1 \leq z\} \times \dots \times \Pr\{X_n \leq z\} \\ &\stackrel{\text{ident. dist.}}{=} \{F(z)\}^n.\end{aligned}$$

**But!** If  $F$  is not known, this is not very helpful because small discrepancies in the estimate of  $F$  can lead to large discrepancies for  $F^n$ .

Need another strategy!

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

#### I. Gumbel

$$G(z) = \exp \left\{ -\exp \left[ -\left( \frac{z-b}{a} \right) \right] \right\}, \quad -\infty < z < \infty$$

#### II. Fréchet

$$G(z) = \begin{cases} 0, & z \leq b, \\ \exp \left\{ -\left( \frac{z-b}{a} \right)^{-\alpha} \right\}, & z > b; \end{cases}$$

#### III. Weibull

$$G(z) = \begin{cases} \exp \left\{ -\left[ -\left( \frac{z-b}{a} \right)^\alpha \right] \right\}, & z < b, \\ 1, & z \geq b \end{cases}$$

with parameters  $a$ ,  $b$  and  $\alpha > 0$ .

## Background on Extreme Value Analysis (EVA)

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### Extremal Types Theorem

If there exist sequences of constants  $\{a_n > 0\}$  and  $\{b_n\}$  such that

$$\Pr \left\{ \frac{M_n - b_n}{a_n} \leq z \right\} \longrightarrow G(z) \text{ as } n \longrightarrow \infty,$$

where  $G$  is a non-degenerate distribution function, then  $G$  belongs to one of the following three types.

## Background on Extreme Value Analysis (EVA)

---

### Extremal Types Theorem

The three types can be written as a single family of distributions, known as the generalized extreme value (GEV) distribution.

$$G(z) = \exp \left\{ - \left[ 1 + \xi \left( \frac{z-\mu}{\sigma} \right) \right]_+^{-1/\xi} \right\},$$

where  $y_+ = \max\{y, 0\}$ ,  $-\infty < \mu, \xi < \infty$  and  $\sigma > 0$ .

## Background on Extreme Value Analysis (EVA)

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### GEV distribution

Three parameters: location ( $\mu$ ), scale ( $\sigma$ ) and shape ( $\xi$ ).

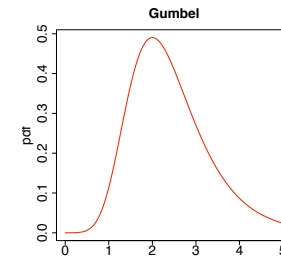
1.  $\xi = 0$  (Gumbel type, limit as  $\xi \rightarrow 0$ )
2.  $\xi > 0$  (Fréchet type)
3.  $\xi < 0$  (Weibull type)

## Background on Extreme Value Analysis (EVA)

---

### Gumbel type

- Light tail
- Domain of attraction for many common distributions (e.g., normal, lognormal, exponential, gamma)

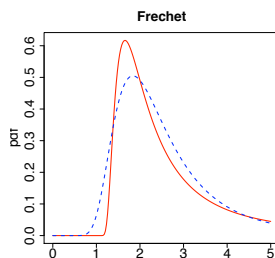


## Background on Extreme Value Analysis (EVA)

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### Fréchet type

- Heavy tail
- $\mathcal{E}[X^r] = \infty$  for  $r \geq 1/\xi$  (i.e., infinite variance if  $\xi \geq 1/2$ )
- Of interest for precipitation, streamflow, economic impacts

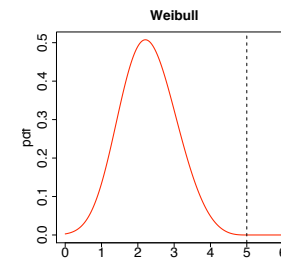


## Background on Extreme Value Analysis (EVA)

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### Weibull type

- Bounded upper tail at  $\mu - \frac{\sigma}{\xi}$
- Of interest for temperature, wind speed, sea level



## Background on Extreme Value Analysis (EVA)

---

### Normal vs. GEV

```
## The probability of exceeding increasingly high values
## (as they double).
# Normal
pnorm( c(1,2,4,8,16,32), lower.tail=FALSE)

# Gumbel
pgev( c(1,2,4,8,16,32), lower.tail=FALSE)

# Fréchet
pgev( c(1,2,4,8,16,32), shape=0.5, lower.tail=FALSE)

# Weibull (note bounded upper tail!)
pgev( c(1,2,4,8,16,32), shape=-0.5, lower.tail=FALSE)
```

## Background on Extreme Value Analysis (EVA)

---

### Normal vs. GEV

```
par( mfrow=c(2,1), mar=c(5,4,0.5,0.5))
plot( 0:20, cdfNorm, ylim=c(0,1), type="l", xaxt="n",
      col="blue", lwd=2, xlab="", ylab="F(x)")
lines( 0:20, cdfGum, col="green", lty=2, lwd=2)
lines( 0:20, cdfFrech, col="red", lwd=2)
lines( 0:20, cdfWeib, col="orange", lwd=2)
legend( 10, 0.05,
       legend=c("Normal", "Gumbel", "Frechet", "Weibull"),
       col=c("blue", "green", "red", "orange"),
       lty=c(1,2,1,1), bty="n", lwd=2)
```

## Background on Extreme Value Analysis (EVA)

---

### Normal vs. GEV

```
# Find  $\Pr\{X < x\}$  for  $X = 0, \dots, 20$ .
cdfNorm <- pnorm( 0:20)
cdfGum <- pgev( 0:20)
cdfFrech <- pgev( 0:20, shape=0.5)
cdfWeib <- pgev( 0:20, shape=-0.5)

# Now find  $\Pr\{X = x\}$  for  $X = 0, \dots, 20$ .
pdfNorm <- dnorm( 0:20)
pdfGum <- dgev( 0:20)
pdfFrech <- dgev( 0:20, shape=0.5)
pdfWeib <- dgev( 0:20, shape=-0.5)
```

## Background on Extreme Value Analysis (EVA)

---

### Normal vs. GEV

```
plot( 0:20, pdfNorm, ylim=c(0,1), typ="l", col="blue",
      lwd=2, xlab="x", ylab="f(x)")
lines( 0:20, pdfGum, col="green", lty=2, lwd=2)
lines( 0:20, pdfFrech, col="red", lwd=2)
lines( 0:20, pdfWeib, col="orange", lwd=2)
```

## Example

---

### Fort Collins, Colorado daily precipitation amount

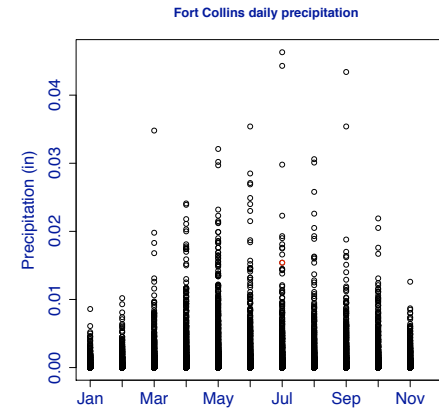
<http://ccc.atmos.colostate.edu/~odie/rain.html>

- Time series of daily precipitation amount (in), 1900–1999.
- Semi-arid region.
- Marked annual cycle in precipitation (wettest in late spring/early summer, driest in winter).
- No obvious long-term trend.
- Recent flood, 28 July 1997. (substantial damage to Colorado State University)

## Example

---

### Fort Collins, Colorado precipitation

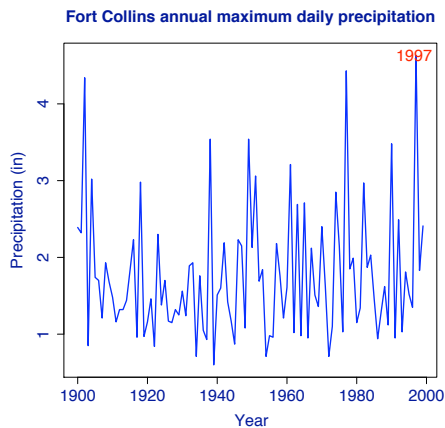


## Example

---

### Fort Collins, Colorado precipitation

#### Annual Maxima



## Example

---

### Fort Collins, Colorado precipitation

How often is such an extreme expected?

- Assume no long-term trend emerges.
- Using annual maxima removes effects of annual trend in analysis.
- Annual Maxima fit to GEV.

```
# Fit Fort Collins annual maximum precipitation to GEV.  
fit <- gev.fit( ftcanmax$Prec/100)
```

```
# Check the quality of the fit.  
gev.diag( fit)
```

## Example

### Fort Collins, Colorado precipitation

Fit looks good (from diagnostic plots).

Parameter	Estimate	(Std. Error)
Location ( $\mu$ )	1.347	(0.617)
Scale ( $\sigma$ )	0.533	(0.488)
Shape ( $\xi$ )	0.174	(0.092)

## Example

### Fort Collins, Colorado precipitation

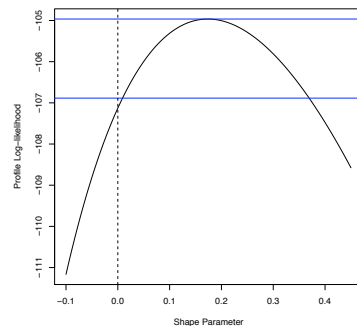
```
# Is the shape parameter really not zero?
# Perform likelihood ratio test against Gumbel type.
fit0 <- gum.fit( ftcanmax$Prec/100)
Dev <- 2*(fit0$nullh - fit$nullh)
pchisq( Dev, 1, lower.tail=FALSE)
```

Likelihood ratio test for  $\xi = 0$  rejects hypothesis of Gumbel type (p-value  $\approx 0.038$ ).

## Example

### Fort Collins, Colorado precipitation

95% Confidence intervals for  $\xi$ , using profile likelihood, are: (0.009, 0.369).



Use `gev.profxi` and `locator(2)` to find CI's.

## Example

### Fort Collins, Colorado precipitation

#### Return Levels

```
# Currently must assign the class "gev.fit" to the
# fitted object, 'fit,' in order to use the "extRemes"
# function, 'return.level'.
class( fit) <- "gev.fit"
fit.rl <- return.level( fit)
```

Return Period	Estimated Return Level (in)	95% CI
10	2.81	(2.41, 3.21)
100	5.10	?(3.35, 6.84)
:	:	:

## Example

---

### Fort Collins, Colorado precipitation

#### Return Levels

CI's from `return.level` are based on the delta method, which assumes normality for the return levels. For longer return periods (e.g., beyond the range of the data), this assumption may not be valid. Can check by looking at the profile likelihood.

```
gev.prof( fit, m=100, xlow=2, xup=8)
```

**Highly skewed!** Using `locator(2)`, a better approximation for the (95%) 100-year return level CI is about (3.9, 8.0).

## Peaks Over Thresholds (POT) Approach

---

Let  $X_1, X_2, \dots$  be an iid sequence of random variables, again with marginal distribution,  $F$ . Interest is now in the conditional probability of  $X$ 's exceeding a certain value, given that  $X$  already exceeds a sufficiently large threshold,  $u$ .

$$\Pr\{X > u + y | X > u\} = \frac{1 - F(u + y)}{1 - F(u)}, y > 0$$

Once again, if we know  $F$ , then the above probability can be computed. Generally not the case in practice, so we turn to a broadly applicable approximation.

## Example

---

### Fort Collins, Colorado precipitation

Probability of annual maximum precipitation at least as large as that during the 28 July 1997 flood (i.e.,  $\Pr\{\max(X) \geq 1.54 \text{ in.}\}$ ).

```
# Using the 'pgev' function from the "evd" package.
pgev( 1.54, loc=fit$mle[1],
      scale=fit$mle[2],
      shape=fit$mle[3],
      lower.tail=FALSE)
```

```
pgev( 4.6, loc=fit$mle[1],
      scale=fit$mle[2],
      shape=fit$mle[3],
      lower.tail=FALSE)
```

## Peaks Over Thresholds (POT) Approach

---

If  $\Pr\{\max\{X_1, \dots, X_n\} \leq z\} \approx G(z)$ , where

$$G(z) = \exp \left\{ - \left[ 1 + \xi \left( \frac{z - \mu}{\sigma} \right) \right]^{-1/\xi} \right\}$$

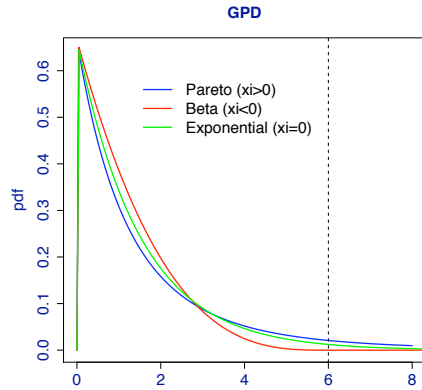
for some  $\mu, \xi$  and  $\sigma > 0$ , then for sufficiently large  $u$ , the distribution  $[X - u | X > u]$ , is approximately the generalized Pareto distribution (GPD). Namely,

$$H(y) = 1 - \left( 1 + \frac{\xi y}{\tilde{\sigma}} \right)_+^{-1/\xi}, y > 0,$$

with  $\tilde{\sigma} = \sigma + \xi(u - \mu)$  ( $\sigma, \xi$  and  $\mu$  as in  $G(z)$  above).

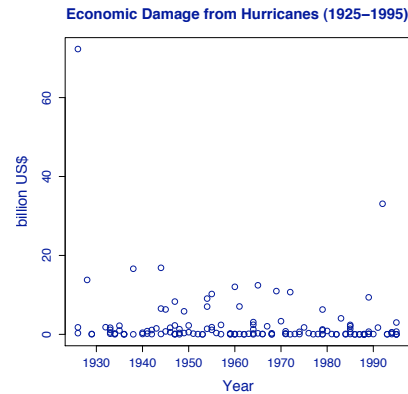
## Peaks Over Thresholds (POT) Approach

- Pareto type ( $\xi > 0$ )  
*heavy tail*
- Beta type ( $\xi < 0$ )  
*bounded above at  $u - \sigma/\xi$*
- Exponential type ( $\xi = 0$ )  
*light tail*



## Peaks Over Thresholds (POT) Approach

### Hurricane damage



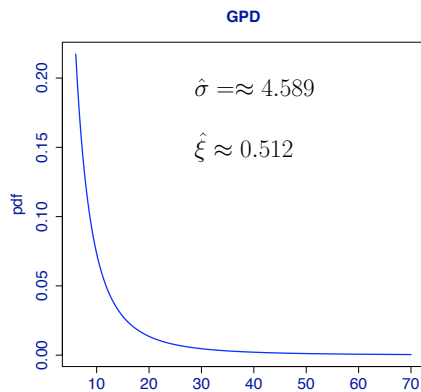
Economic damage caused by hurricanes from 1926 to 1995.

Trends in societal vulnerability removed.

Excess over threshold of  $u = 6$  billion US\$.

## Peaks Over Thresholds (POT) Approach

### Hurricane damage



Likelihood ratio test for

$\xi = 0$  (p-value  $\approx 0.018$ )

95% CI for shape parameter using profile likelihood.  
 $0.05 < \xi < 1.56$

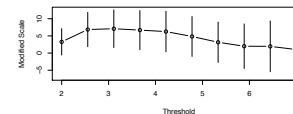
## Peaks Over Thresholds (POT) Approach

### Choosing a threshold

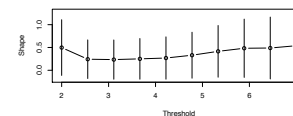
#### Variance/bias trade-off

Low threshold allows for more data (low variance).

Theoretical justification for GPD requires a high threshold (low bias).



`gpd.fitrange( damage$Dam, 2, 7)`



## Peaks Over Thresholds (POT) Approach

---

### Dependence above threshold

Often, threshold excesses are *not* independent. For example, a hot day is likely to be followed by another hot day.

Various procedures to handle dependence.

- Model the dependence.
- De-clustering (e.g., runs de-clustering).
- Resampling to estimate standard errors (avoid tossing out information about extremes).

## Peaks Over Thresholds (POT) Approach

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### Dependence above threshold

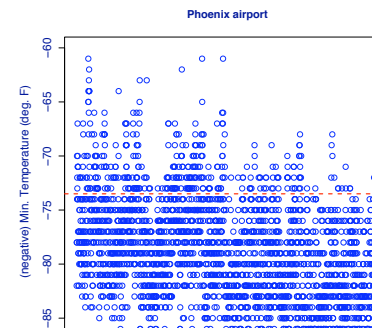
```
# Fit without de-clustering.
phx.fit0 <- gpd.fit( -Tphap$MinT, -73)

# With runs de-clustering (r=1).
phx.dc <- dclust( -Tphap$MinT, u=-73, r=1,
                cluster.by=Tphap$Year)
phxdc.fit0 <- gpd.fit( phx.dc$xdat.dc, -73)
```

## Peaks Over Thresholds (POT) Approach

---

### Dependence above threshold



Phoenix (airport) minimum temperature ( $^{\circ}$ F).

July and August 1948–1990.

Urban heat island (warming trend as cities grow).

Model lower tail as upper tail after negation.

## Peaks Over Thresholds (POT) Approach

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Point Process: *frequency and intensity of threshold excesses*

Event is a threshold excess (i.e.,  $X > u$ ).

Frequency of occurrence of an event (rate parameter),  $\lambda > 0$ .

$\Pr\{\text{no events in } [0, T]\} = e^{-\lambda T}$

Mean number of events in  $[0, T] = \lambda T$ .

GPD for excess over threshold (intensity).

## Peaks Over Thresholds (POT) Approach

---

Point Process: *frequency and intensity of threshold excesses*

Relation of parameters of  $GEV(\mu, \sigma, \xi)$  to parameters of point process  $(\lambda, \sigma^*, \xi)$ .

- Shape parameter,  $\xi$ , identical.
- $\log \lambda = -\frac{1}{\xi} \log \left( 1 + \xi \frac{u - \mu}{\sigma} \right)$
- $\sigma^* = \sigma + \xi(u - \mu)$

More detail: Time scaling constant,  $h$ . For example, for annual maximum of daily data,  $h \approx 1/365.25$ . Change of time scale,  $h'$ , for  $GEV(\mu, \sigma, \xi)$  to  $h'$

$$\sigma' = \sigma \left( \frac{h}{h'} \right)^\xi \text{ and } \mu' = \mu + \frac{1}{\xi} \left\{ \sigma' \left[ 1 - \left( \frac{h}{h'} \right)^{-\xi} \right] \right\}$$

## Peaks Over Thresholds (POT) Approach

---

Point Process: *frequency and intensity of threshold excesses*

Fort Collins, Colorado daily precipitation

Analyze daily data instead of just annual maxima (ignoring annual cycle for now).

### Orthogonal Approach

$$\hat{\lambda} = 365.25 \cdot \frac{\text{No. } X_i > 0.395}{\text{No. } X_i} \approx 10.6 \text{ per year}$$

$$\hat{\sigma}^* \approx 0.323, \hat{\xi} \approx 0.212$$

## Peaks Over Thresholds (POT) Approach

---

Point Process: *frequency and intensity of threshold excesses*

Two ways to estimate PP parameters

- Orthogonal approach (estimate frequency and intensity separately).
  - Convenient to estimate.
  - Difficult to interpret in presence of covariates.
- GEV re-parameterization (estimate both simultaneously).
  - More difficult to estimate.
  - Interpretable even with covariates.

## Peaks Over Thresholds (POT) Approach

---

Point Process: *frequency and intensity of threshold excesses*

Fort Collins, Colorado daily precipitation

Analyze daily data instead of just annual maxima (ignoring annual cycle for now).

### Point Process

$$\hat{\mu} \approx 1.384$$

$$\hat{\sigma} = 0.533$$

$$\hat{\xi} \approx 0.213$$

$$\hat{\lambda} = \left[ 1 + \frac{\hat{\xi}}{\hat{\sigma}} (u - \hat{\mu}) \right]^{-1/\hat{\xi}} \approx 10.6 \text{ per year}$$

## Risk Communication Under Stationarity

---

### Unchanging climate

**Return level**,  $z_p$ , is the value associated with the **return period**,  $1/p$ . That is,  $z_p$  is the level expected to be exceeded on average once every  $1/p$  years.

That is, Return level,  $z_p$ , with  $1/p$ -year return period is

$$z_p = F^{-1}(1 - p).$$

For example,  $p = 0.01$  corresponds to the 100-year return period.

Easy to obtain from GEV and GP distributions (stationary case).

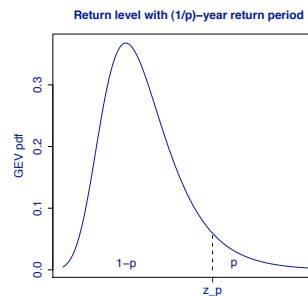
## Risk Communication Under Stationarity

---

### Unchanging climate

For example, GEV return level is given by

$$z_p = \mu - \frac{\sigma}{\xi} [1 - (-\log(1 - p))]^{-\xi}$$



Similar for GPD, but must take  $\lambda$  into account.

## Risk Communication Under Stationarity

---

### Unchanging climate

Compare previous GPD fits (with and without de-clustering). Must first assign the class name, "gpd.fit" to each so that the **extRemes** function, **return.level**, will know whether the objects refer to a GEV or GPD fit.

```
# Without de-clustering.  
class( phx.fit0) <- "gpd.fit"  
return.level( phx.fit0, make.plot=FALSE)
```

```
# With de-clustering.  
class( phxdc.fit0) <- "gpd.fit"  
return.level( phxdc.fit0, make.plot=FALSE)
```

## Non-Stationarity

---

### Sources

- Trends:  
climate change: trends in frequency and intensity of extreme weather events.
- Cycles:  
Annual and/or diurnal cycles often present in meteorological variables.
- Other.

## Non-Stationarity

---

### Theory

No general theory for non-stationary case.

Only limited results under restrictive conditions.

Can introduce covariates in the distribution parameters.

## Non-Stationarity

---

### Phoenix minimum temperature

Recall:  $\min\{X_1, \dots, X_n\} = -\max\{-X_1, \dots, -X_n\}$ .

Assume summer minimum temperature in year  $t = 1, 2, \dots$  has GEV distribution with:

$$\mu(t) = \mu_0 + \mu_1 \cdot t$$

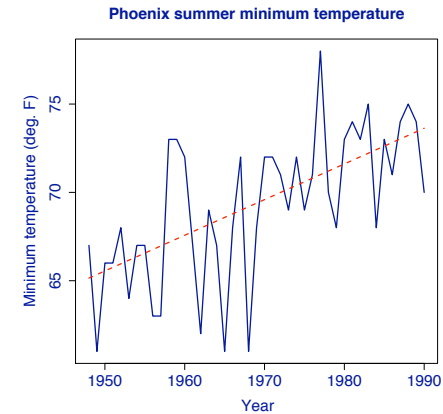
$$\log \sigma(t) = \sigma_0 + \sigma_1 \cdot t$$

$$\xi(t) = \xi$$

## Non-Stationarity

---

### Phoenix minimum temperature



## Non-Stationarity

---

### Phoenix minimum temperature

Note: To convert back to  $\min\{X_1, \dots, X_n\}$ , change sign of location parameters. But note that model is  $\Pr\{-X \leq x\} = \Pr\{X \geq -x\} = 1 - F(-x)$ .

$$\hat{\mu}(t) \approx 66.170 + 0.196t$$

$$\log \hat{\sigma}(t) \approx 1.338 - 0.009t$$

$$\hat{\xi} \approx -0.21$$

Likelihood ratio test

for  $\mu_1 = 0$  (p-value  $< 10^{-5}$ ),

for  $\sigma_1 = 0$  (p-value  $\approx 0.366$ ).

## Non-Stationarity

---

### Phoenix minimum temperature

**Model Checking.** Found the best model from a range of models, but is it a good representation of the data? Transform data to a common distribution, and check the qq-plot.

1. Non-stationary GEV to exponential

$$\varepsilon_t = \left\{ 1 + \frac{\hat{\xi}(t)}{\hat{\sigma}(t)} [X_t - \hat{\mu}(t)] \right\}^{-1/\hat{\xi}(t)}$$

2. Non-stationary GEV to Gumbel (used by `ismev/extRemes`)

$$\varepsilon_t = \frac{1}{\hat{\xi}(t)} \log \left\{ 1 + \hat{\xi}(t) \left( \frac{X_t - \hat{\mu}(t)}{\hat{\sigma}(t)} \right) \right\}$$

## Non-Stationarity

---

### Physically based covariates

Winter maximum daily temperature at Port Jervis, New York

Let  $X_1, \dots, X_n$  be the winter maximum temperatures, and  $Z_1, \dots, Z_n$  the associated Arctic Oscillation (AO) winter index. Given  $Z = z$ , assume conditional distribution of winter maximum temperature is GEV with parameters

$$\mu(z) = \mu_0 + \mu_1 \cdot z$$

$$\log \sigma(z) = \sigma_0 + \sigma_1 \cdot z$$

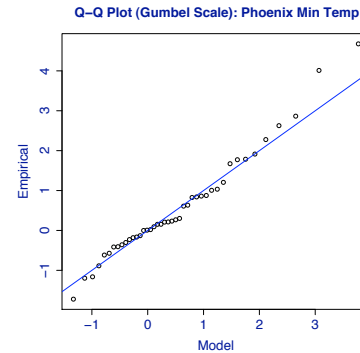
$$\xi(z) = \xi$$

## Non-Stationarity

---

### Phoenix minimum temperature

**Model Checking.** Found the best model from a range of models, but is it a good representation of the data? Transform data to a common distribution, and check the qq-plot.



## Non-Stationarity

---

### Physically based covariates

Winter maximum daily temperature at Port Jervis, New York

$$\hat{\mu}(z) \approx 15.26 + 1.175 \cdot z$$

$$\log \hat{\sigma}(z) = 0.984 - 0.044 \cdot z$$

$$\xi(z) = -0.186$$

Likelihood ratio test for  $\mu_1 = 0$  (p-value  $< 0.001$ )

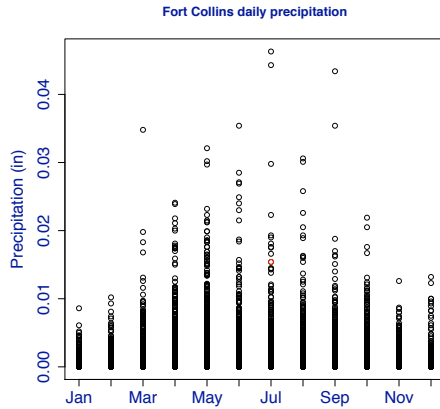
Likelihood ratio test for  $\sigma_1 = 0$  (p-value  $\approx 0.635$ )

## Non-Stationarity

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### Cyclic variation

Fort Collins, Colorado precipitation



## Non-Stationarity

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### Cyclic variation

Fort Collins, Colorado precipitation Orthogonal approach. Next fit GPD with annual cycle in scale parameter.

$$\log \sigma^*(t) = \sigma_0^* + \sigma_1^* \sin\left(\frac{2\pi t}{T}\right) + \sigma_2^* \cos\left(\frac{2\pi t}{T}\right)$$

Giving

$$\log \hat{\sigma}^*(t) \approx -1.24 + 0.09 \sin\left(\frac{2\pi t}{T}\right) - 0.30 \cos\left(\frac{2\pi t}{T}\right)$$

Likelihood ratio test for  $\sigma_1^* = \sigma_2^* = 0$  (p-value  $< 10^{-5}$ )

## Non-Stationarity

---

### Cyclic variation

Fort Collins, Colorado precipitation Orthogonal approach. First fit annual cycle to Poisson rate parameter ( $T = 365.25$ ):

$$\log \lambda(t) = \lambda_0 + \lambda_1 \sin\left(\frac{2\pi t}{T}\right) + \lambda_2 \cos\left(\frac{2\pi t}{T}\right)$$

Giving

$$\log \hat{\lambda}(t) \approx -3.72 + 0.22 \sin\left(\frac{2\pi t}{T}\right) - 0.85 \cos\left(\frac{2\pi t}{T}\right)$$

Likelihood ratio test for  $\lambda_1 = \lambda_2 = 0$  (p-value  $\approx 0$ ).

## Non-Stationarity

---

### Cyclic variation

Fort Collins, Colorado precipitation

Annual cycle in location and scale parameters of the GEV re-parameterization approach point process model with  $t = 1, 2, \dots$ , and  $T = 365.25$ .

$$\mu(t) = \mu_0 + \mu_1 \sin\left(\frac{2\pi t}{T}\right) + \mu_2 \cos\left(\frac{2\pi t}{T}\right)$$

$$\log \sigma(t) = \sigma_0 + \sigma_1 \sin\left(\frac{2\pi t}{T}\right) + \sigma_2 \cos\left(\frac{2\pi t}{T}\right)$$

$$\xi(t) = \xi$$

## Non-Stationarity

---

### Cyclic variation

Fort Collins, Colorado precipitation

$$\hat{\mu}(t) \approx 1.281 - 0.085 \sin\left(\frac{2\pi t}{T}\right) - 0.806 \cos\left(\frac{2\pi t}{T}\right)$$

$$\log \hat{\sigma}(t) \approx -0.847 - 0.123 \sin\left(\frac{2\pi t}{T}\right) - 0.602 \cos\left(\frac{2\pi t}{T}\right)$$

$$\hat{\xi} \approx 0.182$$

Likelihood ratio test for  $\mu_1 = \mu_2 = 0$  (p-value  $\approx 0$ ).

Likelihood ratio test for  $\sigma_1 = \sigma_2 = 0$  (p-value  $\approx 0$ ).

## Risk Communication (Under Non-Stationarity)

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Return period/level does not make sense anymore because of changing distribution (e.g., with time). Often, one uses an "effective" return period/level instead. That is, compute several return levels for varying probabilities over time. Can also determine a single return period/level assuming temporal independence.

$$1 - \frac{1}{m} = \Pr \{ \max(X_1, \dots, X_n) \leq z_m \} \approx \prod_{i=1}^n p_i,$$

where

$$p_i = \begin{cases} 1 - \frac{1}{n} y_i^{-1/\xi_i}, & \text{for } y_i > 0, \\ 1 & \text{, otherwise} \end{cases}$$

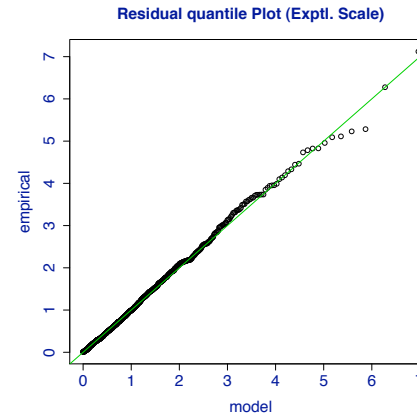
where  $y_i = 1 + \frac{\xi_i}{\sigma_i}(z_m - \mu_i)$ , and  $(\mu_i, \sigma_i, \xi_i)$  are the parameters of the point process model for observation  $i$ . Can be easily solved for  $z_m$  (using numerical methods). Difficulty is in calculating the uncertainty (See Coles, 2001, chapter 7).

## Non-Stationarity

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### Cyclic variation

Fort Collins, Colorado precipitation



## References

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- Coles S, 2001. An introduction to statistical modeling of extreme values. Springer, London. 208 pp.
- Katz RW, MB Parlange, and P Naveau, 2002. Statistics of extremes in hydrology. *Adv. Water Resources*, **25**:1287–1304.
- Stephenson A and E Gilleland, 2006. Software for the analysis of extreme events: The current state and future directions. *Extremes*, **8**:87–109.