

Stochastic 3D Lagrange waves, asymmetry properties and consequences for marine safety

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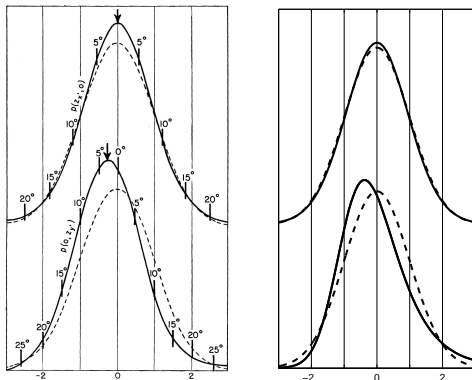
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The Cox and Munk experiment 1954 on wave asymmetry –

compared to Lagrange asymmetry:



Why using a stochastic Lagrange model?

- The (linear) Gaussian wave model allows exact computation of wave characteristic distributions ([the WAFO toolbox](#))
- produces crest-trough and front-back stochastically symmetric waves
- The modified stochastic Lagrange model can produce
 - (2006) crest-trough asymmetric waves (2D)
 - (2009) front-back asymmetric waves (2D)
 - (2011) front-back asymmetric waves (3D)
- still allows for exact computation of wave characteristic distributions!
- Software exists for calculation of statistical wave characteristic distributions (crest, trough heights, period, steepness etc)

Gaussian generator and the orbital spectrum

In the Gaussian model the vertical height $W(t, \mathbf{s})$ of a particle at the free surface at time t and location $\mathbf{s} = (u, v)$ is an integral of harmonics with random phases and amplitudes:

$$\begin{aligned}\boldsymbol{\kappa} &= (\kappa_x, \kappa_y) = \kappa(\cos \theta, \sin \theta) \\ \omega &= \omega(\boldsymbol{\kappa}) = \sqrt{g\kappa \tanh \kappa h} \\ W(t, \mathbf{s}) &= \int_{\omega=0}^{\infty} \int_{\theta=-\pi}^{\pi} e^{i(\boldsymbol{\kappa}\mathbf{s} - \omega t)} d\zeta(\omega, \theta)\end{aligned}$$

with $S(\omega, \theta) =$ the “orbital spectrum” and $\zeta(\omega, \theta)$ is a Gaussian complex “spectral process”.

The stochastic Lagrange model –

Describes vertical and horizontal movements of individual surface water particles. Use

$$W(t, \mathbf{s}) = \int e^{i(\boldsymbol{\kappa}\mathbf{s} - \omega t)} d\zeta(\boldsymbol{\kappa}, \omega)$$

for the vertical movement of a particle with (initial) reference coordinate $\mathbf{s} = (u, v)$ and write

$$\boldsymbol{\Sigma}(t, \mathbf{s}) = \begin{pmatrix} X(t, \mathbf{s}) \\ Y(t, \mathbf{s}) \end{pmatrix} = \text{horizontal location at time } t$$

Make $\boldsymbol{\Sigma}(t, \mathbf{s})$ stochastic.

– with horizontal Gaussian movements

Use the same Gaussian spectral process as in $W(t, \mathbf{s})$ to generate the horizontal variation

Fouques, Krogstad, Myrhaug, Socquet-Juglard (2004), Gjø sund (2000, 2003)
 Aberg, Lindgren, Lindgren (2006, 2007, 2008, 2009, 2011), Guerrin (2009)

$$\boldsymbol{\Sigma}(t, \mathbf{s}) = \begin{pmatrix} X(t, \mathbf{s}) \\ Y(t, \mathbf{s}) \end{pmatrix} = \mathbf{s} + \int \mathbf{H}(\theta, \kappa) e^{i(\kappa \mathbf{s} - \omega t)} d\zeta(\boldsymbol{\kappa}, \omega)$$

where the filter function \mathbf{H} depends on water depth h :

$$\mathbf{H}(\theta, \kappa) = i \frac{\cosh \kappa h}{\sinh \kappa h} \begin{pmatrix} \cos \theta \\ \sin \theta \end{pmatrix}$$

The stochastic Lagrange model

The 3D stochastic first order Lagrange wave model is the triple of Gaussian processes

$$(W(t, \mathbf{s}), \boldsymbol{\Sigma}(t, \mathbf{s})) = (W(t, \mathbf{s}), X(t, \mathbf{s}), Y(t, \mathbf{s}))$$

All covariance functions and auto-spectral and cross-spectral density functions for $\boldsymbol{\Sigma}(t, \mathbf{s})$ follow from the orbital spectrum $S(\omega, \theta)$ and the filter equation.

Front-back asymmetric Lagrange waves

To get realistic front-back asymmetry one needs a model with external input from wind, **for example, by a parameter α** :

$$\frac{\partial^2 X(t, s)}{\partial t^2} = \frac{\partial^2 X_M(t, s)}{\partial t^2} + \alpha W(t, s)$$

The filter function from vertical $W(t, u)$ to horizontal $X(t, u)$ is then

$$H(\omega) = i \frac{\cosh \kappa h}{\sinh \kappa h} + \frac{\alpha}{(-i\omega)^2} = \rho(\omega) e^{i\theta(\omega)},$$

Implies an extra phase shift ($\theta = \pi/2$ in the free model)

$$X(t, u) = u + \int e^{i(\kappa u - \omega t + \theta(\omega))} \rho(\omega), d\zeta(\omega, \kappa)$$

For 3D waves the filter function needs to take directional spreading into account

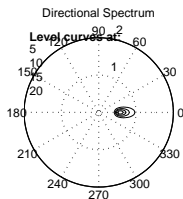
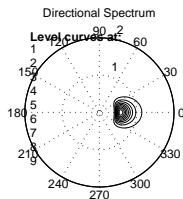
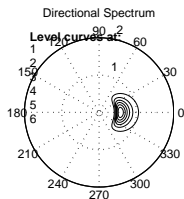
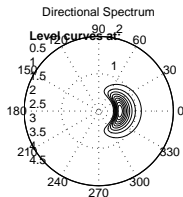
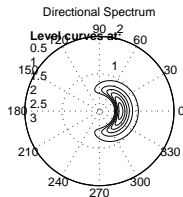
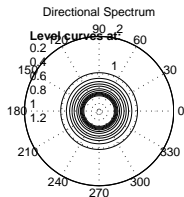
An example:

$$H(\theta, \kappa) = \frac{\alpha}{(i\omega)^2} \cdot \begin{pmatrix} \cos^2(\theta) |\cos(\theta)| \\ \cos^2(\theta) \sin(\theta) \operatorname{sign}(\cos \theta) \end{pmatrix} + i \frac{\cosh \kappa h}{\sinh \kappa h} \cdot \begin{pmatrix} \cos \theta \\ \sin \theta \end{pmatrix}.$$

to take care of wind blowing in positive x-direction.

Front-back asymmetry depends on directional spreading in the orbital spectrum

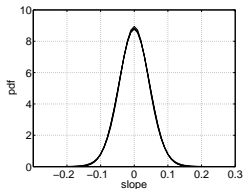
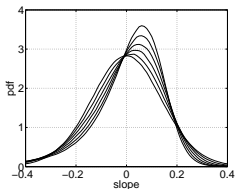
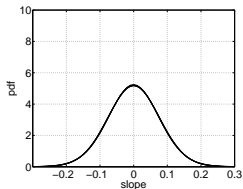
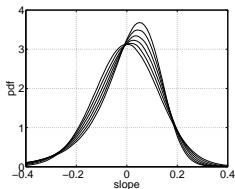
Example spectra - Pierson-Moskowitz with different spreading



Slope PDF across wind and along wind, asynchrone

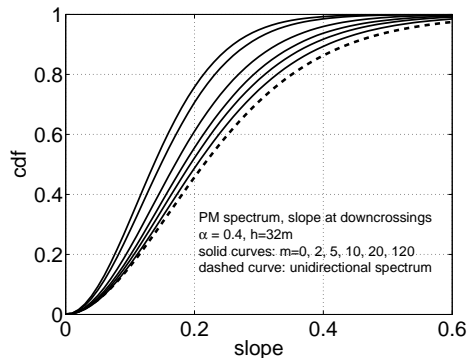
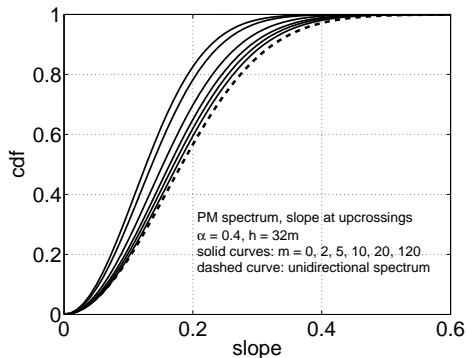
Top: Large spreading – Bottom: Little spreading

Left: Along wind – Right: Across wind; cf. Cox and Munk



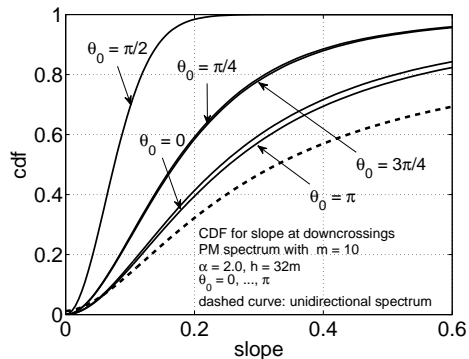
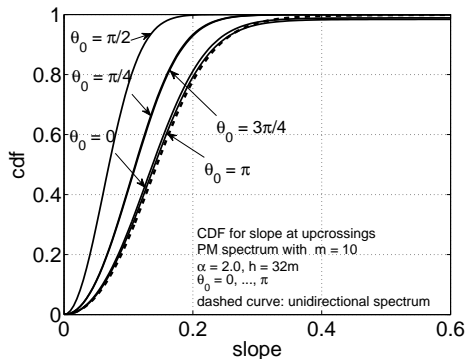
Slope CDF for different directional spreading

Left: slope at upcrossings – Right: Slope CDF at downcrossings
 Moderate linkage parameter: $\alpha = 0.4$



Slope CDF for different heading; Strong linkage: $\alpha = 2$

Left: slope at upcrossings – Right: Slope CDF at downcrossings



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